

# Tracer Advection II: Advanced Numerical Methods for Transport Problems

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## Atmospheric Numerical Modeling : [Desirable Properties]

Numerical algorithms for the next generation atmospheric models should be based on the following criteria:

- Inherent local and global conservation
  - High-order accuracy
  - Computational efficiency
  - Geometric flexibility (complex domain boundaries, AMR capability)
  - Non-oscillatory advection (monotonic or positivity preservation)
  - High parallel efficiency (local method, petascale capability aiming  $O(100K)$  processors)
- 
- Examples of numerical methods which can address the above requirements:-  
Continuous Galerkin or Spectral Element (SE) method, Multimoment Finite-Volume (FV) Method and Discontinuous Galerkin (DG) Method etc..
  - The DG method (DGM) is a hybrid approach which combines nice features of SE and FV methods

## Part-I

- How to solve the basic building block of a complex model – the advection problem – with DGM?

## Flux-Form Atmospheric Equations (Conservation Laws)

- A large class of atmospheric equations of motion for compressible and incompressible flows can be written in **flux (conservation) form**.
- Conservation laws are systems of nonlinear partial differential equations (PDEs) in flux form and can be written:

$$\frac{\partial}{\partial t} U(\mathbf{x}, t) + \sum_{j=1}^3 \frac{\partial}{\partial x_j} F_j(U, \mathbf{x}, t) = S(U),$$

where

- $\mathbf{x}$  is the 3D space coordinate and time  $t > 0$ .  $U(\mathbf{x}, t)$  is the state vector represents mass, momentum and energy etc.
- $F_j(U)$  are given flux vectors and include diffusive and convective effects
- $S(U)$  is the source term
- Linear transport problem is a simple example of conservation law:

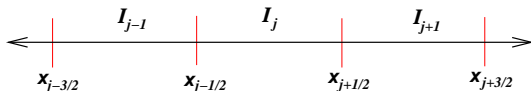
$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \mathbf{V}) = 0, \quad \text{or} \quad \rho_t + \text{div}(\rho \mathbf{V}) = 0$$

# Discontinuous Galerkin Method (DGM) in 1D

- 1D scalar conservation law:

$$\begin{aligned}\frac{\partial U}{\partial t} + \frac{\partial F(U)}{\partial x} &= 0 \quad \text{in } \Omega \times (0, T), \\ U_0(x) &= U(x, t = 0), \quad \forall x \in \Omega\end{aligned}$$

- E.g.,  $F(U) = c U$  (Linear advection),  $F(U) = U^2/2$  (Burgers' Equation)
- The domain  $\Omega$  (periodic) is partitioned into  $N_x$  **non-overlapping** elements (intervals)  
 $I_j = [x_{j-1/2}, x_{j+1/2}]$ ,  $j = 1, \dots, N_x$ , and  $\Delta x_j = (x_{j+1/2} - x_{j-1/2})$



## DGM-1D: Weak Formulation

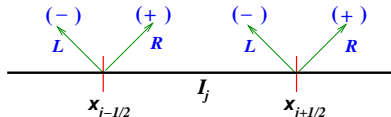
A **weak formulation** of the problem for the approximate solution  $U_h$  is obtained by multiplying the PDE by a **test function**  $\varphi_h(x)$  and integrating over an element  $I_j$ :

$$\int_{I_j} \left[ \frac{\partial U_h}{\partial t} + \frac{\partial F(U_h)}{\partial x} \right] \varphi_h(x) dx = 0, \quad U_h, \varphi_h \in \mathcal{V}_h$$

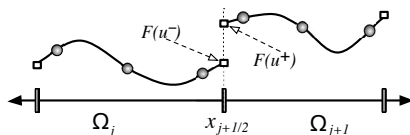
Integrating the second term by parts  $\implies$

$$\int_{I_j} \frac{\partial U_h(x, t)}{\partial t} \varphi_h(x) dx - \int_{I_j} F(U_h(x, t)) \frac{\partial \varphi_h}{\partial x} dx + \\ F(U_h(x_{j+1/2}, t)) \varphi_h(x_{j+1/2}^-) - F(U_h(x_{j-1/2}, t)) \varphi_h(x_{j-1/2}^+) = 0,$$

where  $\varphi(x^-)$  and  $\varphi(x^+)$  denote "left" and "right" limits.



## DGM-1D: Flux term (“Gluing” the discontinuous element edges)



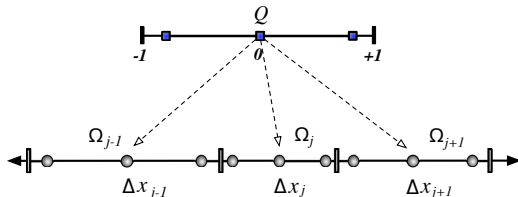
- Flux function  $F(U_h)$  is **discontinuous** at the interfaces  $x_{j\pm 1/2}$
- $F(U_h)$  is replaced by a **numerical flux** function  $\hat{F}(U_h)$ , dependent on the left and right limits of the discontinuous function  $U$ . At the interface  $x_{j+1/2}$ ,

$$\hat{F}(U_h)_{j+1/2}(t) = \hat{F}(U_h(x_{j+1/2}^-, t), U_h(x_{j+1/2}^+, t))$$

- Typical flux formulae (**Approx. Reimann Solvers**): Gudunov, Lax-Friedrichs, Roe, HLLC, etc.
- Lax-Friedrichs numerical flux formula:-

$$\hat{F}(U_h) = \frac{1}{2} \left[ (F(U_h^-) + F(U_h^+)) - \alpha(U_h^+ - U_h^-) \right].$$

## DGM-1D: Space Discretization (Evaluation of the Integrals)



- Map every element  $\Omega_j$  onto the reference element  $[-1, +1]$  by introducing a local coordinate  $\xi \in [-1, +1]$  s.t.,

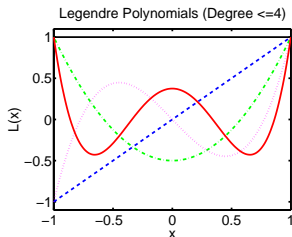
$$\xi = \frac{2(x - x_j)}{\Delta x_j}, \quad x_j = (x_{j-1/2} + x_{j+1/2})/2 \Rightarrow \quad \frac{\partial}{\partial x} = \frac{2}{\Delta x_j} \frac{\partial}{\partial \xi}.$$

- Use a high-order Gaussian quadrature such as the Gauss-Legendre (GL) or Gauss-Lobatto-Legendre (GLL) quadrature rule. The GLL quadrature is 'exact' for polynomials of degree up to  $2N - 1$ .

$$\int_{-1}^1 f(\xi) d\xi \approx \sum_{n=0}^N w_n f(\xi_n); \quad \text{for GLL, } \xi_n \Leftarrow (1 - \xi^2)P'_\ell(\xi) = 0$$



## DGM-1D: Representation of Test function & Approximate Solution



- The **model** basis set for the  $\mathcal{P}^k$  DG method consists of Legendre polynomials,  $\mathcal{B} = \{P_\ell(\xi), \ell = 0, 1, \dots, k\}$ .
- Test function  $\varphi_h(x)$  and approximate solution  $U_h(x)$  belong to  $\mathcal{B}$

$$U_h(\xi, t) = \sum_{\ell=0}^k U_h^\ell(t) P_\ell(\xi) \quad \text{for } -1 \leq \xi \leq 1, \quad \text{where}$$

$$U_h^\ell(t) = \frac{2\ell+1}{2} \int_{-1}^1 U_h(\xi, t) P_\ell(\xi) d\xi \quad \ell = 0, 1, \dots, k.$$

$$\int_{-1}^1 P_m(x) P_n(x) dx = \frac{2}{2m+1} \delta_{m,n} \quad \Leftarrow \text{Orthogonality}$$

- $U_h^\ell(t)$  is the **degrees of freedom (dof)** evolves w.r.t time.

## DGM-1D: Modal Basis Set for a “ $\mathcal{P}^2$ ” Method

- For the  $\mathcal{P}^2$  method,  $\mathcal{B} = \{P_0, P_1, P_2\} = \{1, \xi, (3\xi^2 - 1)/2\}$ .
- Approximate solution:

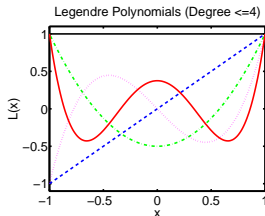
$$U_h(\xi, t) = U_h^0(t) + U_h^1(t) \xi + U_h^2(t) [3\xi^2 - 1]$$

- The *degrees of freedom* to evolve in  $t$  are:

$$U_h^0(t) = \frac{1}{2} \int_{-1}^1 U_h(\xi, t) d\xi \quad \Leftarrow \text{Average}$$

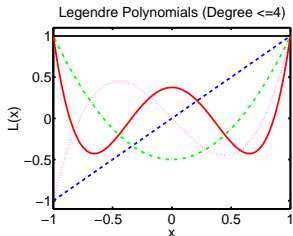
$$U_h^1(t) = \frac{3}{2} \int_{-1}^1 U_h(\xi, t) \xi d\xi$$

$$U_h^2(t) = \frac{5}{2} \int_{-1}^1 U_h(\xi, t) [3\xi^2 - 1] d\xi$$

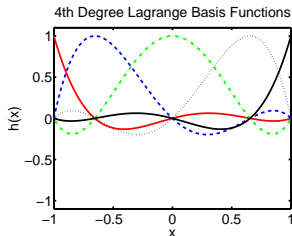


## DGM-1D: Orthogonal Basis Set (Modal Vs Nodal)

Modal basis functions



Nodal basis functions



- The **nodal** basis set  $\mathcal{B}$  is constructed using Lagrange-Legendre polynomials  $h_i(\xi)$  with roots at Gauss-Lobatto quadrature points (**physical space**).

$$U_j(\xi) = \sum_{j=0}^k U_j h_j(\xi) \quad \text{for} \quad -1 \leq \xi \leq 1,$$

$$h_j(\xi) = \frac{(\xi^2 - 1) P'_k(\xi)}{k(k+1) P_k(\xi_j) (\xi - \xi_j)}, \quad \int_{-1}^1 h_i(\xi) h_j(\xi) = w_i \delta_{ij}.$$

- Nodal version was shown to be more computationally efficient than the Modal version (see, *Levy, Nair & Tufo, Comput. & Geos. 2007*)
- Modal version is more “friendly” with monotonic limiting

### DG-1D: Semi-Discretized Form

- Finally, the **weak formulation** leads the PDE to the time dependent ODE

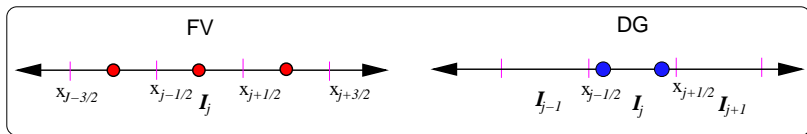
$$\int_{I_i} \left[ \frac{\partial U_h}{\partial t} + \frac{\partial F(U_h)}{\partial x} \right] \varphi_h(x) dx = 0 \Rightarrow \quad \frac{d}{dt} U_h^\ell(t) = \mathcal{L}(U_h) \quad \text{in } (0, T) \times \Omega$$

Example: For the  $\mathcal{P}^1$  case on an element  $I_j$ , we need to solve:

$$\begin{aligned} \frac{d}{dt} U_h^0(t) &= \frac{-1}{\Delta x_j} [F(\xi = 1, t) - F(\xi = -1, t)] \\ \frac{d}{dt} U_h^1(t) &= \frac{-3}{\Delta x_i} \left( [F(\xi = 1, t) + F(\xi = -1, t)] - \int_{-1}^1 U_h(\xi, t) d\xi \right) \end{aligned}$$

Solve the ODEs for the modes at new time level  $U_h^\ell(t + \Delta t)$  For the  $\mathcal{P}^1$  case,

$$U_h(\xi, t + \Delta t) = U_h^0(t + \Delta t) + U_h^1(\xi, t + \Delta t) \xi$$



# Time Integration

- For the ODE of the form,

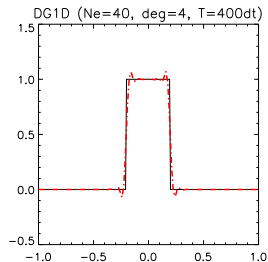
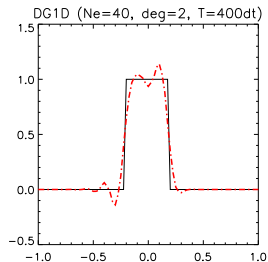
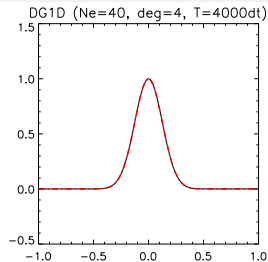
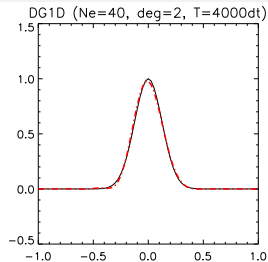
$$\frac{d}{dt} U(t) = \mathcal{L}(U) \quad \text{in } (0, T) \times \Omega$$

- Strong Stability Preserving third-order Runge-Kutta (SSP-RK) scheme (*Gottlieb et al., SIAM Review, 2001*)

$$\begin{aligned} U^{(1)} &= U^n + \Delta t \mathcal{L}(U^n) \\ U^{(2)} &= \frac{3}{4} U^n + \frac{1}{4} U^{(1)} + \frac{1}{4} \Delta t \mathcal{L}(U^{(1)}) \\ U^{n+1} &= \frac{1}{3} U^n + \frac{2}{3} U^{(2)} + \frac{2}{3} \Delta t \mathcal{L}(U^{(2)}). \end{aligned}$$

- CFL for the DG scheme is **estimated** to be  $1/(2k+1)$ , where  $k$  is the degree of the polynomial (*Cockburn and Shu, 1989*).
- Remedy: Use low-order polynomials ( $k \leq 3$ ) or efficient semi-implicit / implicit time integrators or high-order multi-stage R-K method.

## DGM-1D: Results (Simple Linear Advection Test)

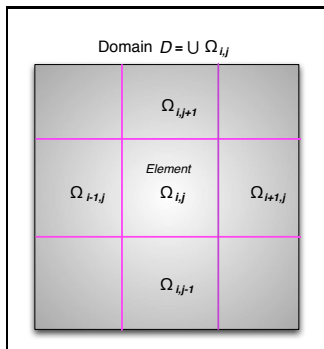


# Discontinuous Galerkin (DG) Methods in 2D Cartesian Geometry

## 2D Scalar conservation law:

$$\frac{\partial U}{\partial t} + \nabla \cdot \mathbf{F}(U) = S(U), \quad \text{in } (0, T) \times \mathcal{D}; \quad \forall (x^1, x^2) \in \mathcal{D},$$

where  $U = U(x^1, x^2, t)$ ,  $\nabla \equiv (\partial/\partial x^1, \partial/\partial x^2)$ ,  $\mathbf{F} = (F, G)$  is the flux function, and  $S$  is the source term.



- The domain  $\mathcal{D}$  is partitioned into non-overlapping elements  $\Omega_{ij}$
- Element edges are discontinuous
- Problem is locally solved on each element  $\Omega_{ij}$

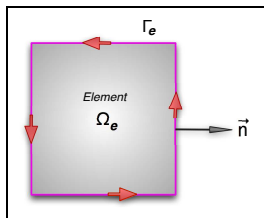
## DG-2D Spatial Discretization for an Element $\Omega_e$ in $\mathcal{D}$

- Approximate solution  $U_h$  belongs to a vector space  $\mathcal{V}_h$  of polynomials  $\mathcal{P}_N(\Omega_e)$ .
- The **Galerkin formulation**: Multiplication of the basic equation by a *test function*  $\varphi_h \in \mathcal{V}_h$  and integration over an element  $\Omega_e$  with boundary  $\Gamma_e$ ,

$$\int_{\Omega_e} \left[ \frac{\partial U_h}{\partial t} + \nabla \cdot \mathbf{F}(U_h) - S(U_h) \right] \varphi_h d\Omega = 0$$

- **Weak Galerkin formulation** : Integration by parts (Green's theorem) yields:

$$\frac{\partial}{\partial t} \int_{\Omega_e} U_h \varphi_h d\Omega - \int_{\Omega_e} \mathbf{F}(U_h) \cdot \nabla \varphi_h d\Omega + \int_{\Gamma_e} \mathbf{F}(U_h) \cdot \vec{n} \varphi_h d\Gamma = \int_{\Omega_e} S(U_h) \varphi_h d\Omega$$



- Orthogonal polynomials (basis functions) are employed for approximating  $U_h$  and  $\varphi_h$  on  $\Omega_e$ .
- Surface and line integrals are evaluated with high-order Gaussian quadrature rule
- Exact Integration: The flux (line) integral should be an order higher than the surface integral (Cockburn & Shu, 1989).

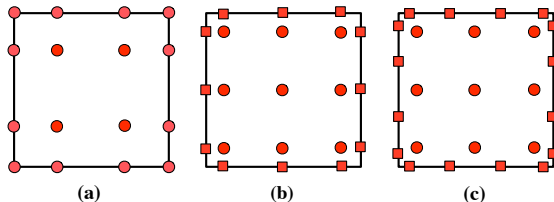


## DG-2D: High-Order Nodal Spatial Discretization

- The nodal basis set is constructed using a tensor-product of Lagrange polynomials  $h_i(\xi)$ , with roots at **Gauss-Lobatto-Legendre** (GLL) or Gauss-Legendre (GL) quadrature points  $\{\xi_i\}$ .

$$[h_i(\xi)]_{GLL} = \frac{(\xi^2 - 1) P'_N(\xi)}{N(N+1) P_N(\xi_i) (\xi - \xi_i)}; \quad \int_{-1}^1 h_i(\xi) h_j(\xi) \simeq w_i \delta_{ij}.$$

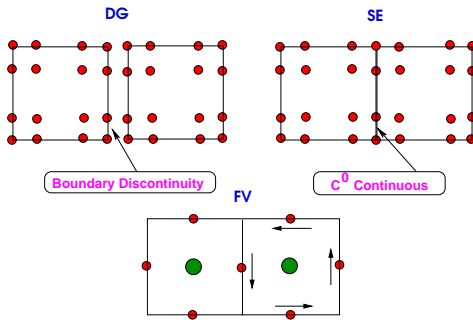
- $P_N(\xi)$  is the  $N^{th}$  degree Legendre polynomial; and  $w_i$  are Gauss quadrature weights



The approximate solution  $U_h$  and test function are represented in terms of nodal basis set.

$$U_{ij}(\xi, \eta) = \sum_{i=0}^N \sum_{j=0}^N U_{ij} h_i(\xi) h_j(\eta) \quad \text{for} \quad -1 \leq \xi, \eta \leq 1,$$

# The DG, SE & FV Methods



- For DGM degrees of freedom (*d.o.f*) to evolve per element is  $N^2$ , where  $N$  is the order of accuracy.
- For FV method the *d.o.f* is 1 (cell-average), irrespective of order of accuracy.
- DGM is based on [conservation laws](#) but exploits the spectral expansion of SE method and treats the element boundaries using FV “tricks.”

# Monotonic Limiter for DG transport

- Importance:

- In atmospheric models, mixing ratios of the advecting chemical species and humidity should be non-negative and free from spurious oscillation.
- The model should avoid creating unphysical negative mass

- Challenges:

- Godunov theorem (1959): “Monotone scheme can be at most first-order accurate”
- There is a “conflict of interest” between the high-order methods and monotonicity preservation!
- In principle, a limiter should eliminate spurious oscillation and preserve high-order nature of the solution to a maximum possible extent

- Existing Limiters for DGM:

- Minmod limiter (Cockburn & Shu, 1989): Based on van Leer's slope limiting, but too diffusive
- Limiters based on WENO or H-WENO (Qui & Shu 2005), Expensive and no positivity preservation
- New bound-preserving limiter: Positivity-preserving and local (Zhang & Shu, 2010)

## Local Bound-Preserving Limiter for DGM

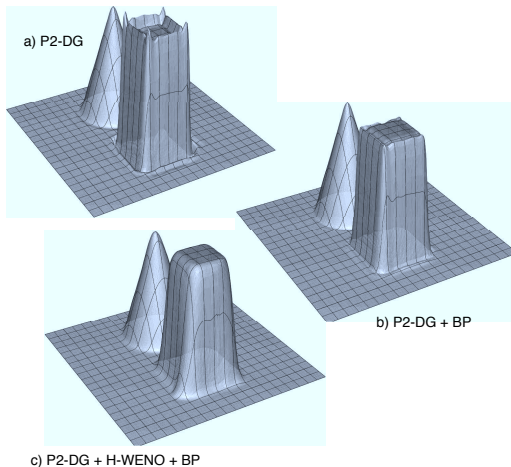
- If the global maximum  $M$  and minimum  $m$  values of the solution  $\rho_{i,j}(x, y)$  is known, then the limited solution  $\tilde{\rho}_{i,j}(x, y)$ :

$$\tilde{\rho}_{i,j}(x, y) = \hat{\theta} \rho_{i,j}(x, y) + (1 - \hat{\theta}) \bar{u}_{i,j}, \quad \hat{\theta} = \min\left\{\left|\frac{M - \bar{u}_{i,j}}{M_{i,j} - \bar{u}_{i,j}}\right|, \left|\frac{m^* - \bar{u}_{i,j}}{m_{i,j}^* - \bar{u}_{i,j}}\right|, 1\right\},$$

- $\bar{u}_{i,j}$  is the average solution in the element  $\Omega_{i,j}$ ,  $M_{i,j} = \max_{(x,y) \in \Omega_{i,j}} \rho_{i,j}(x, y)$  and  $m_{i,j}^* = \min_{(x,y) \in \Omega_{i,j}} \rho_{i,j}(x, y)$ .
- $\hat{\theta} \in [0, 1]$ . The positivity preserving option is a special case of BP filter, and can be achieved by setting  $m^* = 0$ .
- This limiter is conservative and local to the element (Zhang & Shu, JCP, 2010)

## DG Advection on 2D Cartesian Grid (Solid-Body Rotation)

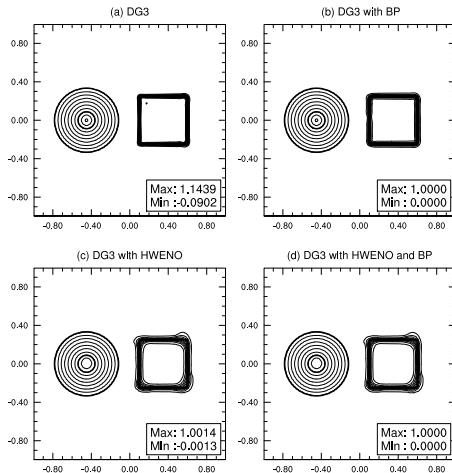
- A DG  $P^2$  (third-order) Model version with 6 DOFs on  $3 \times 3$  G-L grid (Zhang & Nair, MWR, 2012)
- Solid-body rotation (Leveque, 2002) ,  $80 \times 80$  elements.



## DG Advection on 2D Cartesian Grid

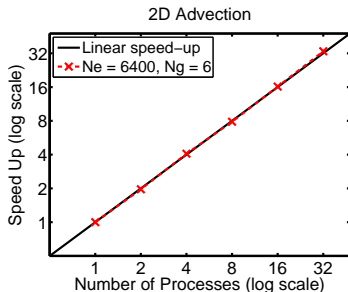
- HWENO uses  $3 \times 3$  cells and completely removes oscillation, but more diffusive.

Leveque data

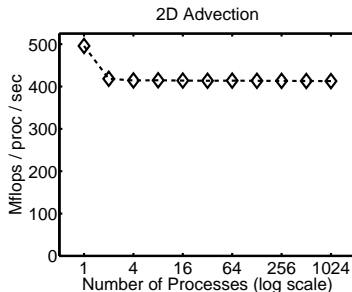


## DG-2D: Scaling Results (Levy, Nair & Tufo, 2007)

- Problem: Advection of a Gaussian-hill,  $80 \times 80$  elements with  $6 \times 6$  GLL grid
- **Strong scaling** is measured by increase the number processes running while keeping the problem size constant
- **Weak scaling** is measured by scaling the problem along with the number of processors, so that work per process is constant



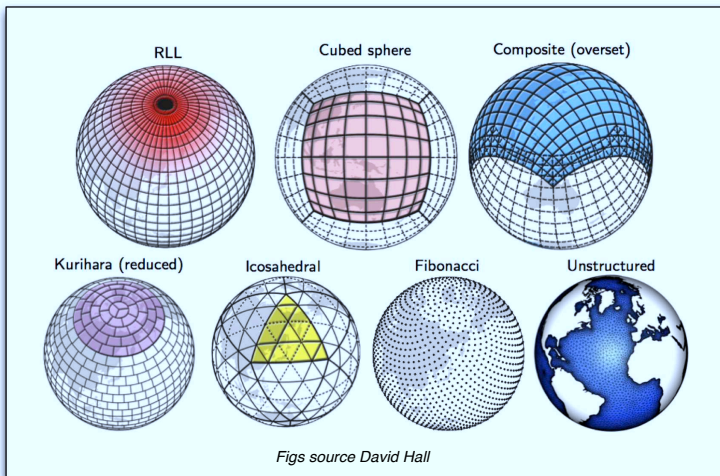
Strong scaling



Weak scaling

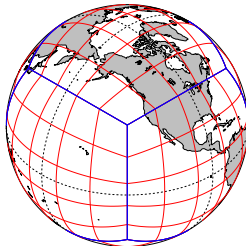
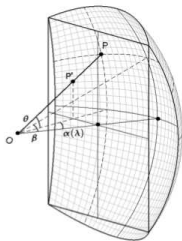
## Extending DG Methods to Spherical Geometry: Various Grid Options

- DG method can be potentially used on various spherical mesh with triangular or quadrilateral (or both) elements





## Cubed-Sphere: Central Equiangular (Gnomonic) Projection



- The sphere is decomposed into 6 identical regions, and free of polar singularities (*Sadourny, MWR, 1972*).
  - Equiangular projection using central angles  $(x^1, x^2)$ .
  - Non-orthogonal grid lines and discontinuous edges
  - All the grid lines are great-circle arcs
  - Quasi-uniform rectangular mesh, well suited for the element-based methods such as DG or SE methods (CAM-HOMME)

# Non-Orthogonal Cubed-Sphere Grid System

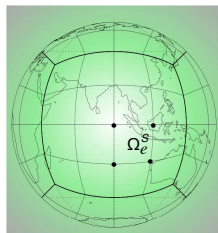
Metric term (Jacobian) of [Cubed-Sphere  $\Rightarrow$  Sphere] Transform on the cubed-sphere:  $\sqrt{G}$

Central angles  $(x^1, x^2) \in [-\pi/4, \pi/4]$ ,  $(\Delta x^1 = \Delta x^2)$  are the independent variables.

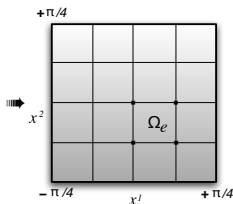
Transport equation (Nair et al. MWR, 2005):

$$\frac{\partial}{\partial t}(\sqrt{G} h) + \frac{\partial}{\partial x^1}(\sqrt{G} u^1 h) + \frac{\partial}{\partial x^2}(\sqrt{G} u^2 h) = 0$$

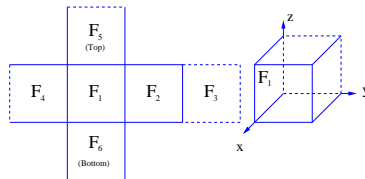
Computational domain is the surface of cube  $[-\pi/4, +\pi/4]^3$



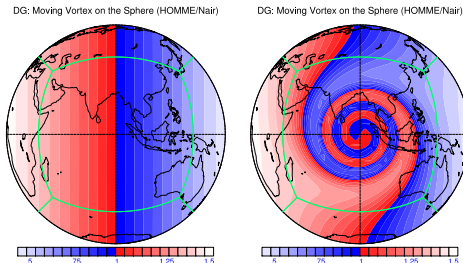
Physical Domain



Computational Domain



## Advection: Deformational Flow (Moving Vortices on the Sphere)



Initial field and DG solution after 12 days. Max error is  $\mathcal{O}(10^{-5})$

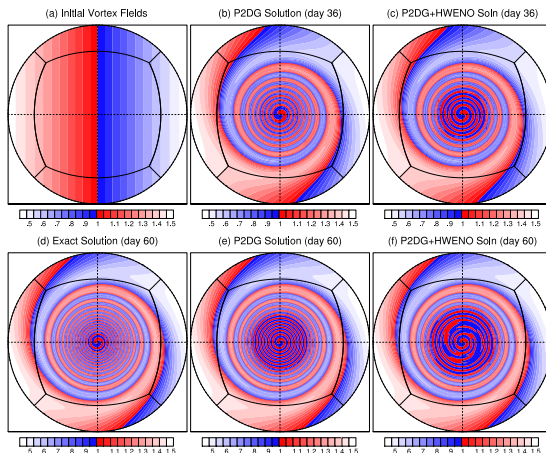
### A Smooth Deformational Flow Test [Nair & Jablonowski (MWR, 2008)]

- The vortices are located at diametrically opposite sides of the sphere, the vortices deform as they move along a prescribed trajectory.
- Analytical solution is known and the trajectory is chosen to be a great circle along the NE direction ( $\alpha = \pi/4$ ).

## DGM Advection: Extreme deformation

- Deformational flow: Fine filament preservation (*Zhang & Nair, MWR, 2012*)
- Modal  $P^2$ -DG with  $100 \times 100 \times 6$  cells,  $\Delta t = 600s$ , 60-day simulation

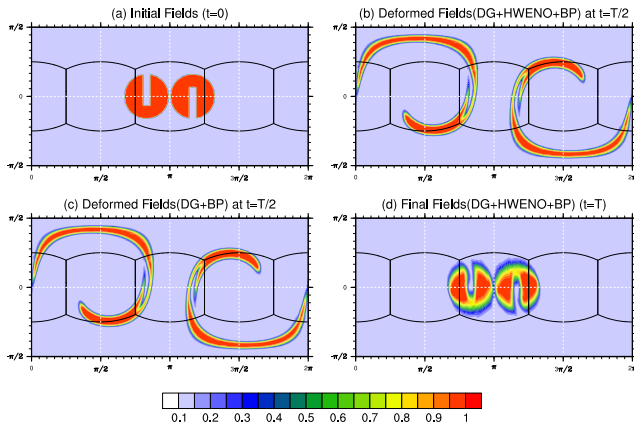
Deformational Flow (Vortex) Test



## DGM Advection: Deformational flow (Slotted-Cylinder)

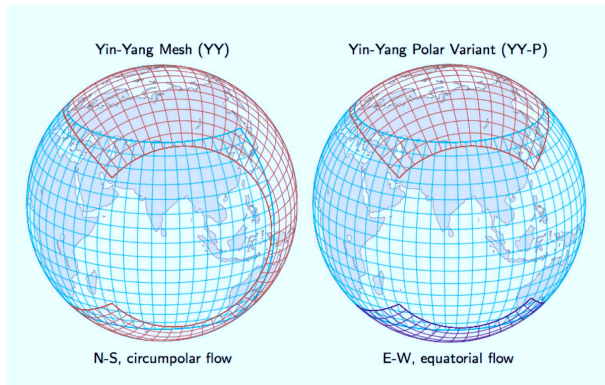
- Deformational flow (non-smooth deformation) (*Nair & Lauritzen, JCP, 2010*)
- Modal  $P^2$ -DG with  $45 \times 45 \times 6$  cells,  $\Delta t = 0.00125s$ ,  $T = 5$ .

Deformational Flow: Slotted Cylinder



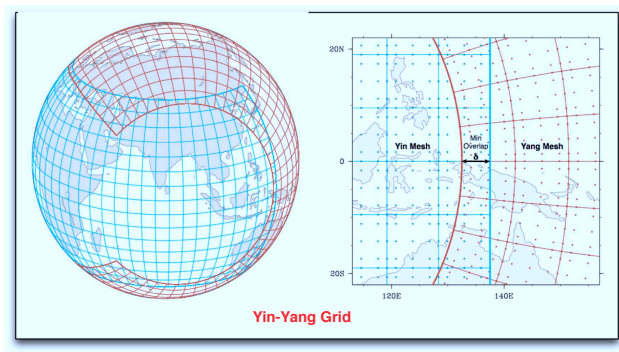
## DG on Yin-Yang Overset Spherical Grid [Kageyama and Sato (2004)]

- It avoids the pole problem of the RLL grid, and there are no singular points.
- The grid spacing is quasi-uniform with a largest to smallest grid-length ratio  $\sqrt{2}$
- Each grid component is orthogonal, producing a simple analytical form for PDEs.
- Overlap regions provide two set of solution.
- Numerical schemes require special treatment for conservation



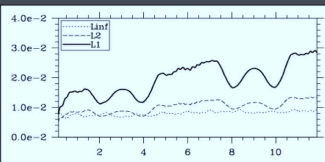
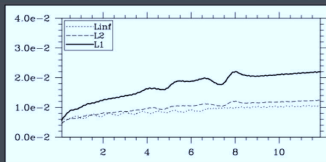
## DG on Yin-Yang Grid: Advection [Hall & Nair, MWR, 2012]

- Sphere  $\mathcal{S} = Y \cup Y'$  where  $Y$ : Yin region, and  $Y'$ : Yang region.  $Y \perp Y'$
- $Y$  is a rectangular region in lat/lon  $(\theta, \lambda)$ -space,  $\lambda \in [-3\pi/2 - \delta, 3\pi/2 + \delta]$ ,  $\theta \in [-\pi/4 - \delta, \pi/4 + \delta]$  where  $\delta$  is the overlap region.
- There are total  $6 \times N_e^2$  elements (DOF) for the DG spatial discretization.

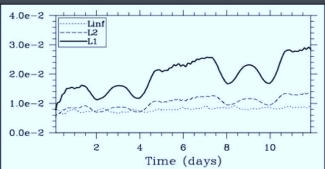
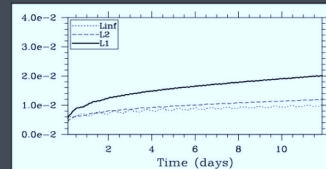


## DG Advection on YY Grid: Cosine-Bell Test (Time-traces of $\ell_1$ , $\ell_2$ , $\ell_\infty$ errors)

Yin-Yang grid at  $\alpha = 0^\circ$  and  $\alpha = 45^\circ$ ,  $N_e = 4$ ,  $N_g = 8$



YY-Polar grid at  $0^\circ$  and  $45^\circ$ ,  $N_e = 4$ ,  $N_g = 8$

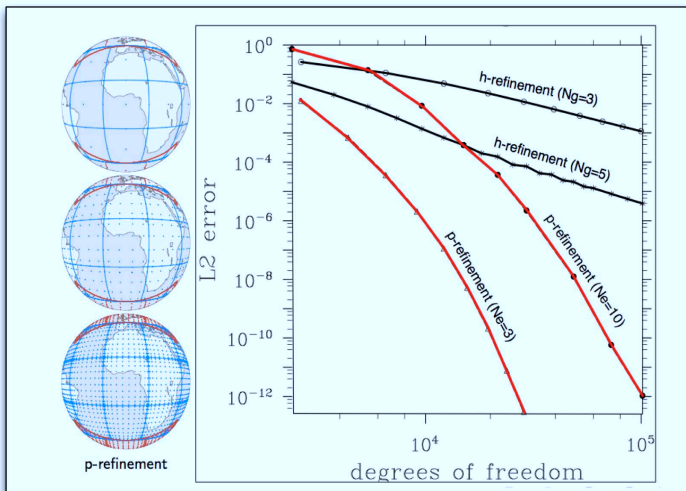


- Cosine bell test results with  $N_e = 4$  and  $N_g = 8$  nodes per element (approximately  $3.2^\circ$  resolution, and 6144 DOF).
- Figs from *Hall & Nair, MWR, 2012*
- Note: Exact mass conservation can be enforced by additional integral constraints (*Baba et al. (2010)*, *Peng et al. (2006)*)



# DGM Convergence: Gaussian Advection and Spectral Convergence

- Advection of a Gaussian Profile (*Levy et al. 2007; Hall & Nair, MWR, 2012*)

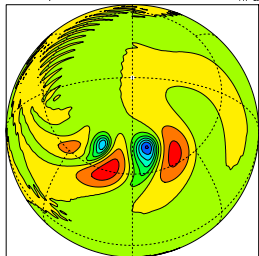


## Beyond Advection: DG-3D Model Vs. CAM Spectral Models

- DG-3D Hydrostatic Dycore (*Nair et al. Comput. & Fluids, 2009*)
- JW-Baroclinic Instability Test, Day 8 Ps ( $\approx 1^\circ$  resolution)
- The DG Solution is smooth and free from “spectral ringing”.

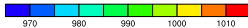
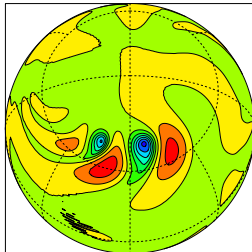
NCAR-T85L26, Day 8

Surface pressure hPa



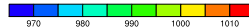
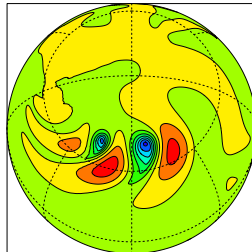
HOMME-SE/Ne30Nv4, Day 8

Surface pressure [res: 1deg] hPa



HOMME-DG/Ne18Nv6, Day 8

Surface pressure [res: 1deg] hPa



### Part-II

- The quasi-Lagrangian coordinates for advection problems

## Hydrostatic Equations in Flux Form: Curvilinear ( $x^1, x^2, \eta$ ) coordinates

$$\frac{\partial u_1}{\partial t} + \nabla_c \cdot \mathbf{E}_1 + \dot{\eta} \frac{\partial u_1}{\partial \eta} = \sqrt{G} u^2 (f + \zeta) - R T \frac{\partial}{\partial x^1} (\ln p)$$

$$\frac{\partial u_2}{\partial t} + \nabla_c \cdot \mathbf{E}_2 + \dot{\eta} \frac{\partial u_2}{\partial \eta} = -\sqrt{G} u^1 (f + \zeta) - R T \frac{\partial}{\partial x^2} (\ln p)$$

$$\frac{\partial}{\partial t} (m) + \nabla_c \cdot (\mathbf{U}^i m) + \frac{\partial(m\dot{\eta})}{\partial \eta} = 0$$

$$\frac{\partial}{\partial t} (m\Theta) + \nabla_c \cdot (\mathbf{U}^i \Theta m) + \frac{\partial(m\dot{\eta} \Theta)}{\partial \eta} = 0$$

$$\frac{\partial}{\partial t} (mq) + \nabla_c \cdot (\mathbf{U}^i q m) + \frac{\partial(m\dot{\eta} q)}{\partial \eta} = 0$$

$$m \equiv \sqrt{G} \frac{\partial p}{\partial \eta}, \nabla_c \equiv \left( \frac{\partial}{\partial x^1}, \frac{\partial}{\partial x^2} \right), \eta = \eta(p, p_s), G = \det(G_{ij}), \frac{\partial \Phi}{\partial \eta} = -\frac{R T}{p} \frac{\partial p}{\partial \eta}.$$

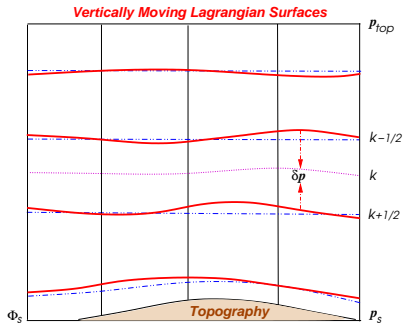
Where  $m$  is the mass function,  $\Theta$  is the potential temperature and  $q$  is the moisture variable.  $\mathbf{U}^i = (u^1, u^2)$ ,  $\mathbf{E}_1 = (E, 0)$ ,  $\mathbf{E}_2 = (0, E)$ ;  $E = \Phi + \frac{1}{2} (u_1 u^1 + u_2 u^2)$  is the energy term.  $\Phi$  is the geopotential,  $\zeta$  is the relative vorticity, and  $f$  is the Coriolis term.

[Ref: HOMME/DG, Nair et al. *Comput. & Fluids* 2009]

## Vertical (quasi) Lagrangian Coordinates (*Starr, J. Meteorol. 1945*)

A “vanishing trick” for vertical advection terms!

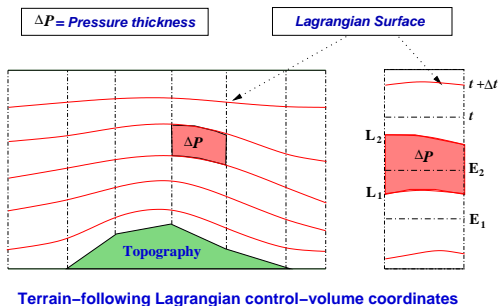
- Terrain-following Eulerian surfaces are treated as material surfaces ( $\dot{\eta} = 0$ ).
- Simplified hydrostatic equations with no “vertical terms”
- The resulting **Lagrangian surfaces** are free to move up or down direction.



# The Remapping of Lagrangian Variables

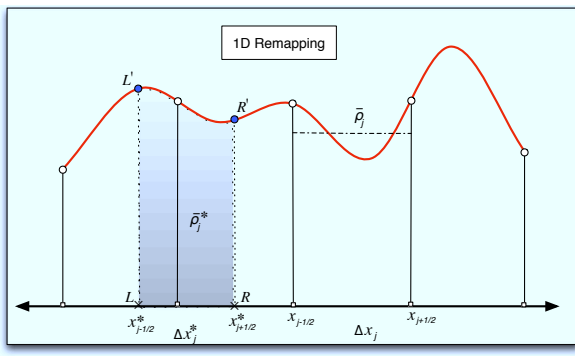
## Vertically moving Lagrangian Surfaces

- Over time, Lagrangian surfaces deform and must be remapped.
- The velocity fields ( $u_1, u_2$ ), and total energy ( $\Gamma_E$ ) are **remapped** onto the reference coordinates.



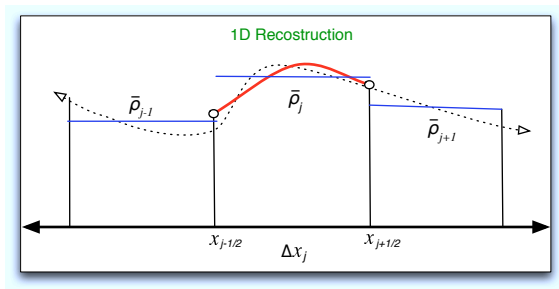
Remapping: Lauritzen & Nair, MWR, 2008; Norman & Nair, MWR, 2008)

## Remapping (Rezoning or Re-gridding) on a 1D Grid



- Remapping: Interpolation from a source grid to target grid with constraints (conservation, monotonicity, positivity-preservation etc.).
- Application: Conservative semi-Lagrangian methods (e.g. CSLAM); Grid-to-grid data transfer for pre- or post-processing (GeCore).

## Remapping (Rezoning or Re-gridding) on a 1D Grid



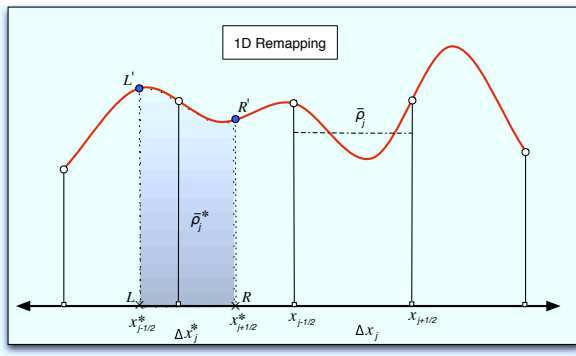
- **Reconstruction:** Fit a piecewise polynomial  $\rho_j(x)$  for every cell  $\Delta x_j = x_{j+1/2} - x_{j-1/2}$ , using the known cell-average values  $\bar{\rho}_j$  from the neighboring cells.
- The subgrid-scale distribution  $\rho_j(x)$  must satisfy the conservation constraint:

$$\bar{\rho}_j = \frac{1}{\Delta x_j} \int_{x_{j-1/2}}^{x_{j+1/2}} \rho_j(x) dx, \quad \Rightarrow \text{Mass} = \bar{\rho}_j \Delta x_j$$

- $\rho_j(x)$  may be further modified to be monotonic (E.g: PLM, PPM, PCM, PHM)



## Remapping 1D Grid



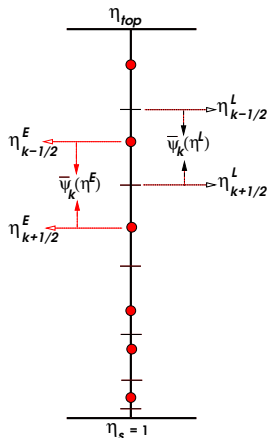
- Mass in the target cell ( $\Delta x_j^* = x_{j+1/2}^* - x_{j-1/2}^*$ ) can be expressed as the difference of "Accumulated Mass" ( $\mathcal{A}_m$ ):

$$\bar{\rho}_j^* \Delta x_j^* = \mathcal{A}_m(RR') - \mathcal{A}_m(LL') \Rightarrow \bar{\rho}_j^* = \frac{1}{\Delta x_j^*} [\mathcal{A}_m(RR') - \mathcal{A}_m(LL')]$$

where

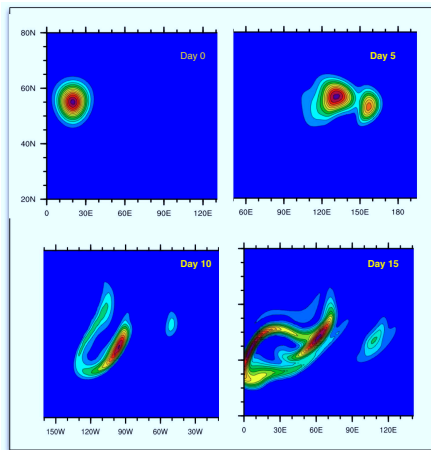
$$\mathcal{A}_m(RR') = \int_{x_{Ref}}^{x_{j+1/2}^*} \rho(x) dx = \sum_{k=1}^{j-1} \bar{\rho}_k \Delta x_k + \int_{x_{j-3/2}}^{x_{j+1/2}^*} \rho_{j-1}(x) dx$$

# Vertical Advection with Lagrangian $\eta$ -Coordinate



- Reference (initial) grid  $\eta = \eta(P, P_s) \in [\eta_{top}, 1]$
- Source grid = Lagrangian  $\eta_k^L$
- Target grid = Eulerian  $\eta_k^E$ ;  $\sum \Delta \eta_k^E = \sum \Delta \eta_k^L$
- Lagrangian  $\eta_k^L$  can be computed from the predicted “pressure thickness”  $\Delta P$  (CAM-FV)
- Remapping is performed at every advective  $\Delta t$
- Every 1D vertical trajectory information can be “recycled” for all tracers

## 3D Transport ( CAM-SE): SE horizontal + vertical remapping



- CAM-SE ( $1^\circ$ ): JW-Test divergent flow field. SE horizontal transport is quasi-monotonic
- $\Delta t_a = 4 \times 90$  s, vertical remapping by PCM (Zerroukat, 2005) for advection.

Figure courtesy: Christoph Erath

## Summary & Conclusions :

- The DG method with moderate order (third or fourth) is an excellent choice for transport problems as applied in atmospheric sciences. DGM addresses:
  - 1 High-order accuracy
  - 2 Geometric flexibility
  - 3 Positivity-preserving advection
  - 4 High parallel efficiency
  - 5 Local and global conservation
- In comparison with finite-volume and finite-difference implementations of the Yin-Yang grid, the DG approach is considerably simpler as the overset interpolation is local, requiring information from the interior of a single element.
- In general, modified YY-P and YY meshes exhibited similar performance on most tests, while the YY-P mesh performed better on cases with strictly zonal flow.
- DG method is an ideal candidate for the new generation petascale-capable dynamical cores.
- The “moving” vertical Lagrangian (evolve and remap approach) method provides an efficient way for 3D conservative multi-tracer transport.

THANK YOU!